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On the two-day oscillations and the day-to-day variability in global 3-D-modeling of the chemical system of the upper mesosphere/mesopause region

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Abstract. The integration of the photochemical system of the upper mesosphere/mesopause region brought evidence that the system is able to respond in a nonlinear manner under certain conditions. Under the action of the diurnally-periodic insolation, the system creates subharmonic oscillations or chaos if disregarding strong diffusion, and under special conditions it possesses multiple solutions. The models used in the past were simplified and idealized in view of the number of dimensions and the consideration of the full dynamics. On the basis of our global 3-D-model of the dynamics and chemistry of the middle atmosphere (COMMA-IAP), we also found a nonlinear response in the photochemistry under realistic conditions. The model under consideration is not yet self-consistent, but the chemical model uses the dynamical fields calculated by the dynamic model. From our calculations we got period-2 oscillations of the photochemical system within confined latitudinal regions around the solstices but not during the equinoxes. The consequence of the period-2 oscillation of the chemical active minor constituents is that a marked two-day variation of the chemical heating rates is an important thermal pumping mechanism. We discuss these findings particularly in terms of the influence of realistic dynamics on the creation of nonlinear effects.

1 Introduction

As early as 1987 Sonnemann and Fichtelmann found that the photochemical system of the mesosphere represents a nonlinear enforced chemical oscillator driven by the diurnally-periodic insolation. This system is able to produce chemical resonances if the characteristic chemical response time ranges in the order of 1 day. Precisely this is true in the mesopause region. It was shown in the joint paper by Fichtel-

mann and Sonnemann (1987) that this system is able to create subharmonic oscillations of period-2 and period-3 and chaos depending on the period of an artificial day of insolation. We call such phenomena nonlinear effects. The period of a subharmonic oscillation – in short subharmonics – is a multiple of the fundamental period amounting to a day in the case of the chemistry of the earth atmosphere which is driven by the diurnally periodic solar insolation. In an ideal case for a period- n , every n days the same diurnal variations for all constituents repeat whereas the diurnal variations of the individual days differ within this period of n days. For $n \rightarrow \infty$ a so-called deterministic chaos arises. As early as 1918 Duffing showed that such periodically driven nonlinear oscillators were able to create cascades of subharmonic oscillations. There is a rather large body of publications dealing with this subject, and we refer here only to a standard work such as that of Thompson and Stewart (1991) describing its history and fundamentals or the excellent review by Feigin (2002) who summarized and analyzed the current state of the art regarding nonlinear processes in atmospheric photochemical systems. Abarbanel (1995) treated the problem of the analysis of observed chaotic data comprehensively and thoroughly. The same subject applied to the atmospheric chemistry considering observed chaotic time series has also been explored by Feigin et al. (2001). Some theoretical investigations and methodical problems of the nonlinear response of the tropospheric photochemical system have been examined, among others, in publications by Krol and Poppe (1998) and Poppe (1999).

The highly idealized system of Sonnemann and Fichtelmann (1987), considering only the odd oxygen constituents O and O₃ and the odd hydrogen species H, OH and HO₂ as variable compounds (phase variables), was improved later on. The behavior of this dynamical zero-dimensional model was systematically investigated using the water vapor concentration, the ratio of day-time to nighttime hours (reflecting the season and/or latitude) and the air density (reflecting the geometric height)

as control parameters (Fichtelmann and Sonnemann, 1992; Sonnemann and Fichtelmann, 1997). Particularly the paper by Sonnemann and Fichtelmann (1997) investigated the behavior from point of view of chaos research. Yang and Brasseur (1994) showed that the photochemical system of the mesosphere possesses a trigger solution if taking into calculation fluxes from the thermosphere into the mesosphere. In the meantime the Russian group from Nizhny Novgorod further analyzed the nonlinear response of this system and derived the so-called essential model of the chemistry of the mesopause region (Konovalov et al., 1997; Feigin et al., 1998; Konovalov and Feigin, 2000; Feigin et al., 2001, Feigin, 2002). They found that the effect results in essence from an interplay between atomic oxygen and atomic hydrogen. The nonlinear behavior of the system was further investigated by Montecinos (1996); Johnson et al. (1998), Montecinos and Felmer (1999) and Scott et al. (2000). Montecinos and Felmer (1999) found that the period-2 is a stable oscillation of the photochemical system. They also reported about multiple solutions.

The control parameter “frequency of solar insolation” seems to be an unrealistic parameter for the earth’s atmosphere as the earth rotation period of 24 h is constant. However, it was shown by Sonnemann (2001) that the zonal wind moving an air parcel around a parallel results in a Doppler shift of the period of solar insolation. Under real conditions the shift of the period amounts to between +4 and −7 h. On the basis of a simplified 1-D-model defined around a parallel and considering a changing zonal wind only, the system created subharmonics again when the zonal wind exceeded a critical velocity. This so-called photochemical Doppler effect has been studied in more detail on the basis of three-dimensional calculations by Sonnemann and Grygalashvyly (2003), but a nonlinear response was not observed.

Kulikov and Feigin (2004) investigated reaction-diffusion waves in the mesospheric photochemical system caused by horizontal eddy diffusion. It was shown that when the spectrum of oscillation contains a marked second subharmonics, reaction-diffusion waves in the form of traveling fronts and pulses of the phase of two-day oscillations may be excited.

The influence of the turbulent diffusion on the nonlinear behavior was investigated by means of a one-dimensional model (Sonnemann and Feigin, 1999a and 1999b; Sonnemann et al., 1999) using the diffusion coefficient as a control parameter. A very complicated system response has been found. Different subtle and catastrophic so-called space-bifurcations creating subharmonic and chaotic attractors occur for a changing diffusion coefficient. The nonlinear response seems to be restricted to a certain height interval, but actually, the differences between the concentration amplitudes for consecutive days becomes smaller the more distant the height from this interval is. This behavior has been called the onion bifurcation. The most essential finding was, however, that strong diffusion suppresses the nonlinear response of the system. The critical values of the diffusion coefficient for the creation of a first subharmonic oscilla-

tion by a subtle bifurcation, a period-2 oscillation, ranges in the order of $K \approx 2 \times 10^5 \text{ cm}^2 \text{ s}^{-1}$ comparable with real relations at the lower and middle mesopause region. As frequently very small diffusion coefficients of the order of $K \approx n \times 10^4 \text{ cm}^2 \text{ s}^{-1}$ have been derived, particularly around and below 80 km (e.g. Lübken 1997), the question arose whether a nonlinear response of the photochemical system can also occur under real conditions. In other words, in this paper we will not only contribute to the basic research in nonlinear dynamics, but we will go a step toward a more realistic description of the complicated behavior of the chemistry in the mesopause region. We will search for conditions under which subharmonics either could really occur or, on the other hand, conditions under which an appearance cannot be expected or rather it is very unlikely. We will also point out the shortcomings of the model under consideration and state what we have to improve in order to get reliable results.

2 Brief model description

For calculations we use our global three-dimensional model COMMA-IAP of the dynamics and chemistry of the middle atmosphere (0–150 km). The acronym stands for Cologne Model of the Middle Atmosphere of the Institute of Atmospheric Physics in Kühlungsborn. It was described in more detail e.g. in Berger (1994); Sonnemann et al. (1998); Kremp et al. (1999); Berger and von Zahn (1999); Körner and Sonnemann (2001); Sonnemann and Körner (2003), Hartogh et al. (2004). The model is designed to investigate particularly the extended mesopause region and the mesosphere. It considers all relevant species and chemical families. The resolution amounts to about 1.1 km in altitude, 5 (or 2.5) deg in latitude and 5.625 deg in longitude. For the following calculations we start with a time step of 900 s but change the variable time resolution to 1 min after 1 year of model time. At the same time we integrate O and O₃ no longer only as family, but separately. The system then operates close to a flowing equilibrium but the change to the time-consuming integration represents a certain perturbation of the system so that the transients normally die away sufficiently only after about one or two weeks in that altitude. A flowing equilibrium expresses the fact that the concentration of short-lived constituents closely follows the slowly changing system conditions such as solar insolation, temperature or the concentration of long-lived source gases. It does not, of course, represent a steady state. The chemical transport model uses the dynamical fields of the temperature, the wind components and the pressure from the dynamical model. It considers the advective transport and both the turbulent and molecular diffusion. The model calculates climatological means. It is unable to create such phenomena as planetary waves or stratospheric warmings. The model does not operate interactively. The chemical model uses the dynamical fields; yet the dynamical model is not fed online with data from the chemical model.

The quality of the model used has been shown in various papers dealing with very different scientific subjects. Kremp et al. (1999) showed, in comparison with radar wind observations, that the seasonal wind patterns are correctly calculated, and in the article by Berger and von Zahn (1999) even the two level mesopause was modeled correctly. A touchstone for the correctness of the important vertical wind is the calculation of the seasonal variation of the middle atmospheric water vapor concentration. Körner and Sonnemann (2001) and, using an improved model version, Sonnemann and Grygalashvily (2005) computed, among other things, the seasonal behavior of the water vapor concentration. Particularly the results in the second paper mirror the seasonal variation observed, and even such special phenomenon as the areas of the possible occurrence of noctilucent clouds (NLC) are sufficiently well calculated. In Hartogh et al. (2004) the so-called tertiary ozone maximum close to the polar night terminator was correctly reproduced.

As the diffusion possesses the ability to destroy nonlinear effects, the treatment of all diffusive processes is an important methodical problem. The turbulent or eddy diffusion coefficient reflects chaotic processes of the air (for example breaking of gravity waves) which cannot be resolved in the frame of the dynamic model. The eddy diffusion coefficient is an external parameter for the model. There is a large uncertainty and lack of observations relating to the spatio-temporal behavior of the eddy diffusion coefficient in the mesopause region. Actually, it depends on height, latitude, season and local time. There are different estimations (e.g. Hocking 1990; Lübken 1997, etc.) and indirect hints (e.g. from meteor trail observations) on the magnitude of the diffusion coefficient showing a large scatter over two orders in magnitude in the mesopause region. For calculations we use our so-called standard eddy diffusion profile – an idealized model with a diffusion coefficient of $K=2\times 10^5 \text{ cm}^2 \text{ s}^{-1}$ at 80 km increasing to a maximum value of $5.1\times 10^5 \text{ cm}^2 \text{ s}^{-1}$ at 90 km. These values exceed even that value of the first creation of a period-2 in case of the 1-D-model (Sonnemann and Feigin, 1999a).

In view of the creation of nonlinear effects in the photochemistry, we are interested in the consideration of relatively quiet conditions. However, very important is the suppression of the numerical diffusion due to the treatment of the vertical advection. The vertical numerical diffusion coefficient is given by $D_{num}=0.5*(\Delta z*|w|-\Delta t*w^2)$ (e.g. Smolarkiewicz 1984) with w -vertical velocity and Δz -height resolution and Δt -time resolution. The second term is very small, but the first term has an order of $n\times 10^5 \text{ cm}^2 \text{ s}^{-1}$ and larger within the mesopause region comparable with or even stronger than the eddy diffusion coefficient. Particularly the strong vertical tidal winds result in a marked numerical diffusion whereas the zonally (diurnally) averaged vertical winds are one order in magnitude smaller than the tidal winds and possess values of maximum few cm s^{-1} .

The procedure of the so-called anti-numerical diffusion reduces the uncertainties resulting from the advective transport scheme of a constituent marked by a certain concentration

gradient which tends to flatten this gradient (for an extended discussion see e.g. Prather 1986 or Walcek, 2000). There are different procedures of anti-numerical diffusion used (e.g. Smolarkiewicz, 1983; Prather, 1986; Bott, 1989; Bott, 1992) which reduced the diffusivity of the transport scheme but they did not suppress it sufficiently well, particularly under conditions of steep concentration gradients. In our opinion, this fact could be one of the main reasons that global models have not produced a nonlinear response thus far.

However, we cannot assess the models developed by other groups, as information about the numerical treatment is normally outlined only very scanty and limited. Generally, the numerical diffusion works in addition to the turbulent diffusion so that an effective acting diffusion coefficient can be defined as the sum of both parts. A large diffusivity influences the distribution of the minor constituents such as water vapor just within the mesopause region. So the water vapor mixing ratio in the area of NLC at 83 km at high latitudes in summer calculated by employing the Smolarkiewicz scheme amounted to about 4 ppmb (Körner and Sonnemann, 2001).

Now, we have implemented a new advective transport scheme developed by Walcek and Aleksic (1998) and Walcek (2000) which is almost free (nearly zero) of numerical diffusion. The water vapor mixing ratio increased to about 6 ppmv after the change and now agrees very well with measurements by HALOE of the Upper Atmospheric Research Satellite (UARS) displaying all typical latitudinally-seasonal patterns (Sonnemann and Grygalashvily, 2004; Sonnemann et al., 2004).

The water vapor concentration is, apart from the diffusion coefficient and the vertical wind, the most important control parameter for the consideration of nonlinear effects. The height region of possible nonlinear response was perhaps too dry in the older model versions to create nonlinear effects, but more likely the system was too diffusive and the time resolution too poor. The dissociation rate of water vapor occurs only combined with the water vapor concentration in the chemical code. It determines the hydrogen radical production term of the chemical scheme. The dissociation rate of water vapor is chiefly determined by the solar Lyman- α flux above about 75 km. Compared with the paper by Körner and Sonnemann (2001), the dissociation rates have been corrected according to the flux values given by Woods et al. (2000).

The new measurements by UARS increased the old flux values by about 55 %. For calculations here we use a Lyman- α flux according to the year 2001 being close to solar maximum conditions. The changes in the model - the smaller time step, the strong reduction of the numerical diffusion also entailing higher water vapor mixing ratios within the mesopause region and the increase of the water vapor dissociation rate – promote the creation of nonlinear effects.

3 Problems of the integration of 3-D-models

There were some hints of a nonlinear response of the complex global three-dimensional model in the past but they had not been considered in more detail thus far. A general drawback of three-dimensional models is connected with the family concept of the integration of the chemical code. Due to the extremely large computer time needed to integrate a very rigid system, an integration of all individual constituents separately is impossible. Their individual characteristic chemical times extend over many orders in magnitude. Particularly, the odd hydrogen and odd oxygen species have, depending on height, extremely small characteristic times. Special attention must be paid to the times of sunrise and sunset because the solar insolation changes greatly during a relatively short time period and the system comes out of its almost flowing equilibrium. However, within a chemical family the individual members will only be transformed into each other without changing their family concentration. The family concentration is conserved. An example for this assertion are the reactions $\text{H} + \text{O}_3 \rightarrow \text{OH} + \text{O}_2$ and $\text{OH} + \text{O} \rightarrow \text{H} + \text{O}_2$ with the net reaction $\text{O} + \text{O}_3 \rightarrow 2\text{O}_2$. The odd oxygen species are catalytically destroyed but the concentration of the odd hydrogen constituents remains constant in this cycle. Hence, the characteristic family time (defined in Sonnemann and Fichtelmann, 1997) is essentially greater than those of the individual family members. Consequently, the families are subjected to transport. The characteristic family times of both the odd families ranges in the order of 1 day around 80 to 85 km whereas the individual characteristic chemical times possess minimum values in the order of seconds up to minutes. We also have to note that in case of a nonlinear response of the system, the response time of the entire system can be essentially greater than the family times if the system operates in the vicinity of a bifurcation point. This can create long-lasting transients of any subharmonics or chaos. A crucial problem is also the fact that the characteristic times drastically change from the lower to the upper model domain and between daytime and the night when no photodissociation acts.

In order to integrate the entire system within an acceptable computer time, in three-dimensional models the necessary integration time step is often in the order or even essentially larger than the characteristic times of the acting constituents. The integration concept used distinguishes between such species which can be immediately put into a (flowing) equilibrium without having to be time-integrated and transported (e.g. $\text{O}(^1\text{D})$ and other excited species or the plasma constituents within the D- or E-layer), species which have to be time-integrated, and such species which can be derived from the family concentration. The characteristic transport times range in the order of days or longer in the middle atmosphere (see also Körner and Sonnemann, 2001). The system considers such constituents which have to be time-integrated and transported (such as water vapor, methane, dinitrogen oxide, etc.) and the odd families (such as HO_x or NO_x), which also have to be time-integrated and transported.

A serious problem, as mentioned above, consists in the partition of the family concentration into those of the individual members. A necessary requirement is, of course, the conservation of the number of atoms. This is also a request for all procedures of the anti-numerical diffusion. Under this condition the concentration of the individual members can be calculated on the basis of their current equilibrium concentrations. This procedure loses its validity if a short-living constituent becomes a long-living constituent at a certain height, time of day, season or latitude. This is, for instance, the case for atomic oxygen and atomic hydrogen above the mesopause or for ozone in lower altitudes and during the night etc. Thus we integrate atomic oxygen and ozone separately but this line of action requires essentially to reduce the time step.

Under the condition that the concentrations of a constituent varies close to a flowing equilibrium, the integration time step can be essentially enlarged compared with the necessary time step for conditions far from a flowing equilibrium. Particularly around sunset and sunrise, the dissociation rates vary strongest with time and the chemical system is substantially disturbed. The chemical system runs behind the permanently changing radiation conditions. The characteristic time of this perturbation could also be defined by that time during which the dissociation rate changes (decreases or increases) by a factor of $e=2718$ related to the noon value. There is no step function for the diurnal variation of the dissociation rates in reality as used in the older investigations (e.g. Fichtelmann and Sonnemann, 1987, 1992 and Sonnemann and Fichtelmann, 1997) but, according to the real sunrise and sunset, a more or less smoothed variation of their values depending on latitude, season and species. Consequently, we use dissociation rates according to solar zenith angles as far as 95° in steps of 1 degree. The characteristic time ranges in the order of roughly an hour or larger and consequently, the integration time step should be noticeably below that time for the fastest constituents, which have to be time integrated. In order to be on the safe side, we use a time step of 1 min. A reduction to 15 s yielded virtually the same results and this was also true for an enlarged time step of 150 s. In contrast, in case of non-separated treatment of the odd oxygen constituents we used a time step from 450 s to 5400 s (see also Körner, 2002).

We are aware how difficult a trustworthy calculation of a nonlinear response on the basis of such large and complex system is. It was shown by means of the simplified models that even a small change of the control parameters can result in a catastrophic or subtle bifurcation. A subharmonics of a certain period changes into another period or into chaos if the conditions are altered only a little. Then the integration time step is rather relevant for the concrete system response. A strong restriction also results from the coarse space resolution particularly with regard to the height resolution of the order of 1 km. The simplified models have vertical resolutions in the order of 100 m and sometimes below that. However, when considering diffusion, the period-2 oscillation is the most prominent subharmonics and the first

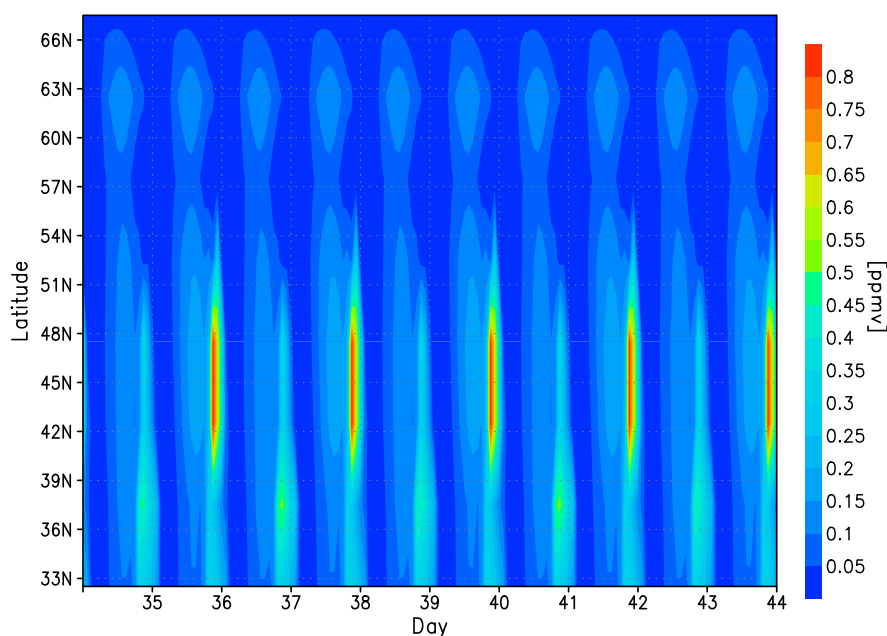


Fig. 1. Ozone mixing ratio of the Northern Hemisphere at 83 km altitude for a permanent 1st July showing a latitudinal region of period-2 oscillation between about 33° N and 56° N. The transients nearly died away.

nonlinear response of the system when it comes from a limit cycle of one day period and exceeds a critical control parameter value. The bifurcation is, as mentioned, subtle in this case. The period-2 occupied the largest control parameter range (such as water vapor concentration, air density, length of the daytime, diffusion coefficient, etc.) from the entire range of nonlinear response. In other words, a period-2 variation as usually results from the calculations is a very robust subharmonics.

A serious problem in view of observations also consists in the difficulty in deciding between chemically induced subharmonics or even chaos and dynamically imprinted variations. Idealized models can calculate the system response under fixed conditions, such as a constant date and constant aeronomic conditions, varying only one control parameter separately. Under real conditions all parameters change permanently and thus, we have to expect only episodes of a certain system behavior (e.g. the occurrence of a certain subharmonic oscillation within a limited space of time) within certain geographic regions. We start the search for a nonlinear response of the chemical transport model with fixed dynamical and seasonal conditions according to a permanent fixed date. In this case the model runs until this date with a coarse time resolution (one year, in order to come close to a flowing equilibrium). After that we use the same diurnal variation of all dynamical parameters and dissociation rates according to this date. Then we integrate the system with these data for a long period using a small time step now and transport O and O₃ separately. In the second case – the so-called real case – we proceed in the same manner but after a given date we change only the time step and integration procedure but do not fix the seasonal conditions.

4 Results

Figure 1 shows a survey of calculations of the ozone mixing ratio of the Northern Hemisphere at 83 km altitude for a permanent 1st July after the transients died away (after 34 days). Then the system operates in an almost limit cycle mode. The tick numbers stand for the days left of the tick marks. We show the ozone mixing ratios being indicative for the other constituents because ozone is the best measurable constituent in the mesopause region. The figure clearly displays a period-2 oscillation from about 33° N to 56° N. The largest amplitude difference occurs around 46° N. Around 37.5° N a weak period-4 occurs. Only the fundamental period has been found on the Southern Hemisphere up to the low northern latitudes. Figure 2 displays, according to the results depicted in Fig. 1, an altitude-time section at 42.5° N of the same constituent. The figure makes clear that the two-day oscillation occurs approximately between 80 and 85 km with a maximum close to 83 km. Figure 3 displays the same state of affairs at 72.5° N outside of the occurrence range of period-2 oscillation shown in Fig. 1. There is practically no period-2 at 83 km but a pronounced two-day oscillation between 86 and 90 km with a maximum between 88 and 89 km. A phase jump occurs between 87 and 88 km. Generally in our model, the height region of the nonlinear response increases with increasing latitude but strongest for the highest latitudes. The same calculation carried out for the equinoxes does not yield a strong indication of a period-2 oscillation anywhere.

The calculations carried out for the south summer season do not result in a simply mirrored behavior. This finding results primarily from the dynamic model which is based on the

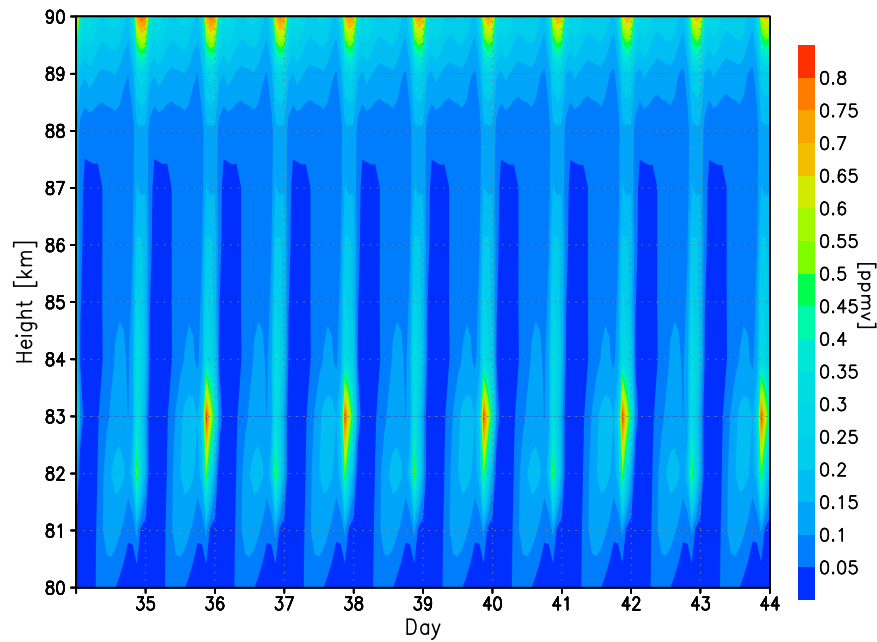


Fig. 2. Altitude-time section at 42.5° N according to the results presented in Fig. 1. The figure demonstrates that the domain of nonlinear response is confined to between about 80 and 85 km.

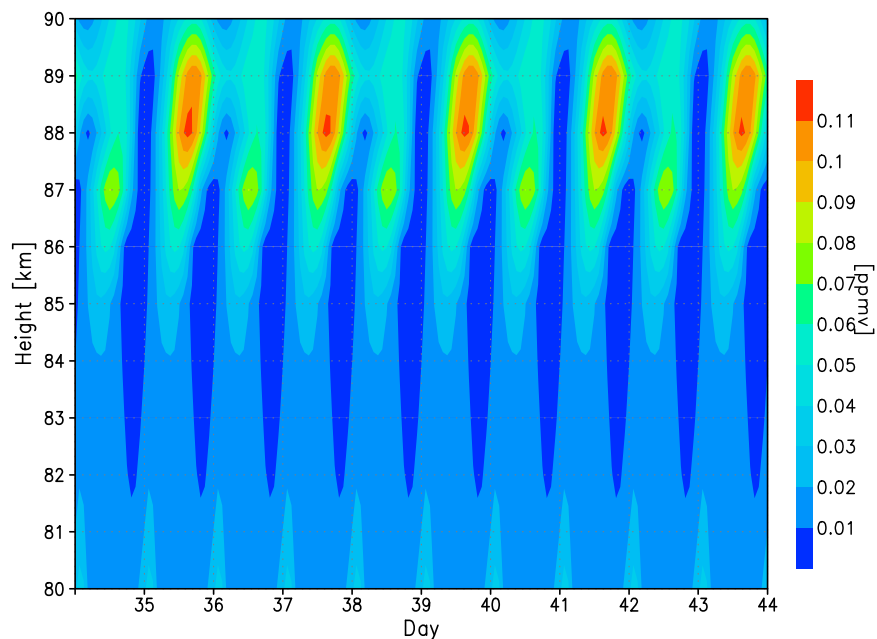


Fig. 3. The same state of affairs as shown in Fig. 2 but at 72.5° N. Now the region of nonlinear response occurs between 86 and 90 km whereas at 83 km no period-2 was found (compare with Fig. 1).

Berlin ozone climatology of the stratosphere (Pawson et al., 1998) being asymmetrical with respect to the hemispheres under comparable seasonal conditions and it is partly due to the eccentricity of the earth's orbit. Figure 4 shows a cutout between 14° and 36° S of the ozone mixing ratio at 83 km for a permanent 22nd December. Now the effect occurs in the southern hemisphere in middle to low latitudes but also in middle to low latitudes of the Northern Hemisphere. The po-

sition of the regions of nonlinear response changes with the date. Period-2 oscillations occur during the entire summer season again. The figure displays clear period-2 oscillations, however, with a phase shift between about 25° S and 26° S.

The period-2 is the preferred subharmonics, but as Fig. 5 demonstrates the next period doubling can also be created by the model. The figure shows the diurnal variations of ozone at 2.5° N at 83 km height for a permanent 31st July after a

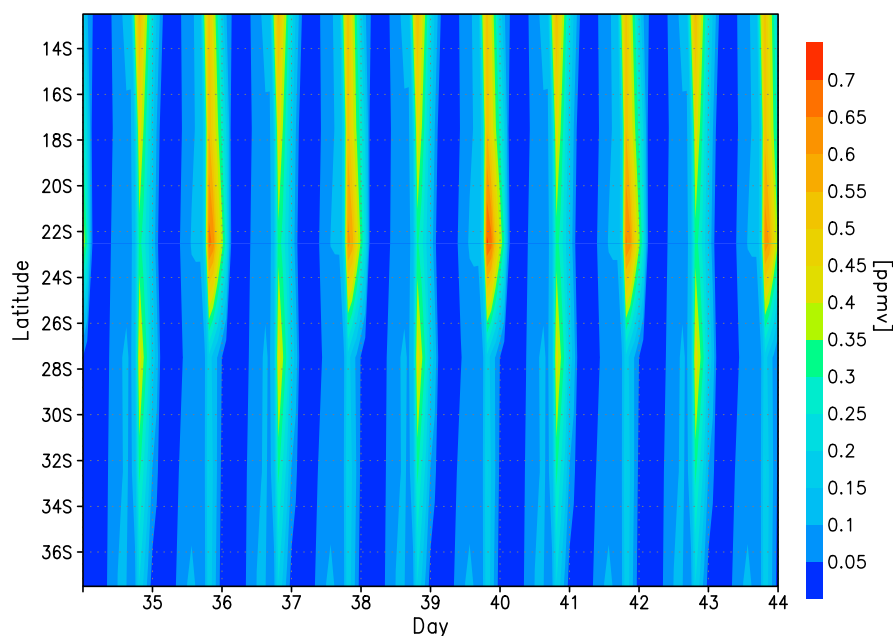


Fig. 4. A cutout between 14° and 36° S of the ozone mixing ratio at 83 km for a permanent 22nd December. There is a phase shift with respect to the maximum between about 25° S and 26° S.

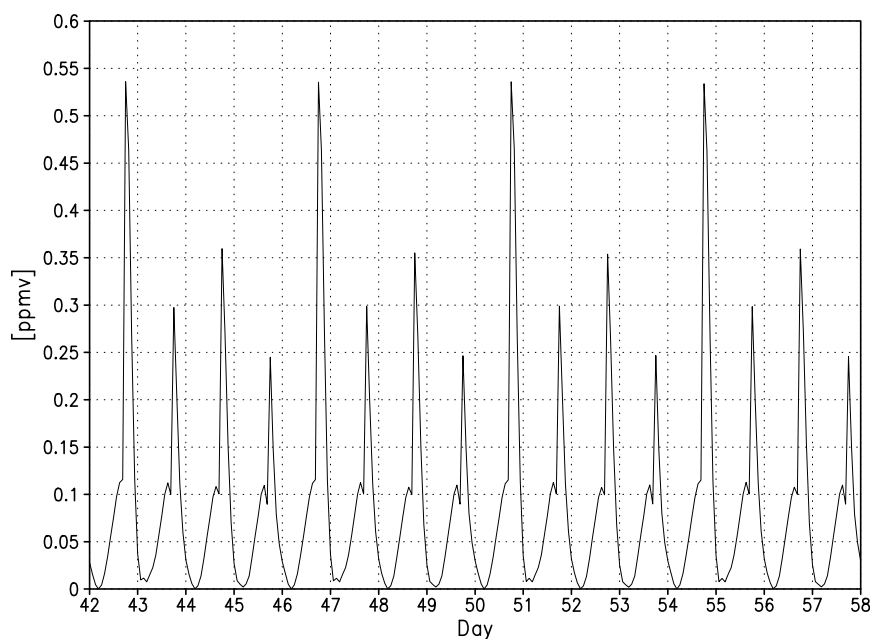


Fig. 5. Diurnal variations of ozone at 2.5° N at 83 km height for a permanent 31st July after a long time of calculation showing a clear period-4 oscillation, which demonstrates the potential of the system to create a period doubling cascade.

long time of calculation. The system operates nearly in a limit cycle mode. In reality such amplitude differences are masked by the natural variability and one may observe an alternation between large and small amplitudes. Figure 6 shows an example at 83 km and 12.5° N for a permanent 31st January. Now the representation starts instantly after the change of the solution procedure. Evidently, the system responds immediately with the creation of a transient period-

2 oscillation. After some periods a transient quasi period-6 occurs and that changes finally into a period-2 again approximating nearly a limit cycle.

Figure 7 depicts the diurnal variations of the total chemical heating rates for 4 consecutive days at different altitudes and latitudes for the 1st July run. It is evident that the two-day oscillation of the individual species is connected with a marked thermal two-day variation of the chemical heating. Figure 8

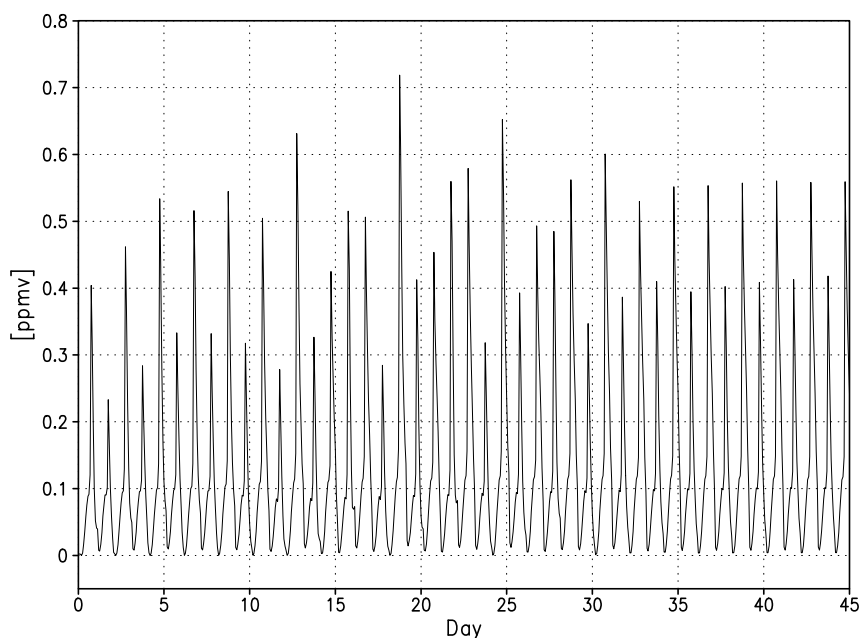


Fig. 6. Example of the occurrence of transient subharmonics at 83 km and 12.5° N for a permanent 31st January. The figure shows results instantly after the change of the integration procedure. A transient period-2 changes into a transient period-6 and that returns to a period-2 again.

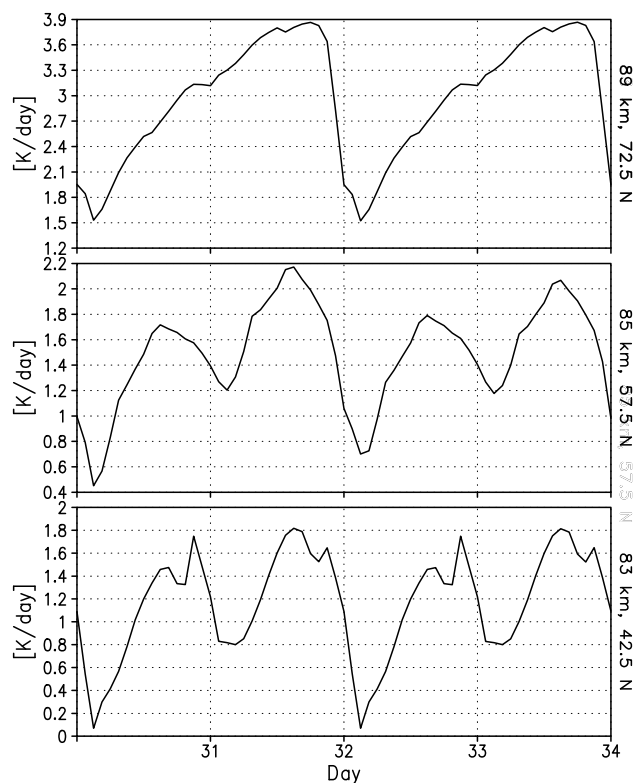


Fig. 7. Diurnal variations of the total chemical heating rate for 4 consecutive days for a permanent 1st July run at different altitudes and latitudes.

displays the diurnal variations of the 4 most important chemically active species for four consecutive days at 62.5° N and at 84 km for a permanent 1st July. All constituents show clear two-day variations of their mixing ratio.

A trajectory describes the movement of a phase point within a sub-phase space and an attractor reflects the motion in case of a limit cycle. Attractors are indicative for the non-linear response of the system. Figure 9 displays an example of a quasi attractor for ozone versus atomic oxygen at 83 km height and 42.5° N for a permanent 1st July. The data of calculations were stored every 1.5 h and therefore the curves appear to be very angular. In reality they are smoothed. Although the calculations based on a permanent date the figure do not represent an ideal limit cycle because the long-living constituents still change slowly during the time period considered. Both the different diurnal variations can be clearly recognized as separate loops.

The results presented here clearly show that a so-called “two-day wave” in the chemistry of the mesopause region can occur under favorable conditions but that the two-day variation is not a general feature of the chemical dynamics. In order to use defined conditions, we integrated the system under the conditions of a continuously lasting date. In reality the conditions change slightly according to the running date. This concerns the slightly seasonally changing radiation conditions but it refers more to all influences connected with the changing dynamics. As in climatological mean the dynamical patterns does not normally drastically change within a period of one or two weeks, we also expect under real conditions the creation of, at least, a transient period-2 oscillation in favorable locations. Figure 10 demonstrates this statement

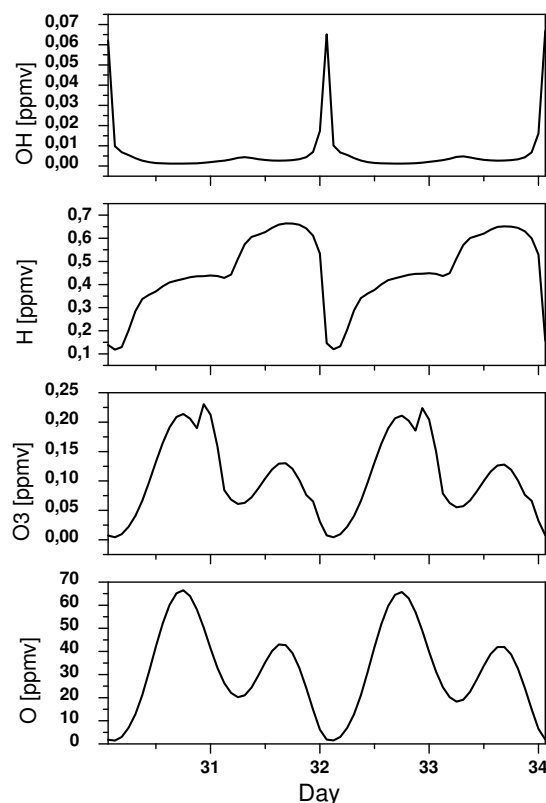


Fig. 8. Diurnal variations of the 4 most important chemically active species (odd oxygen O, O₃ and odd hydrogen H, OH constituents) for four consecutive days at 84 km and 62.5° N for a permanent 1st July.

for the north summer season. The figure shows a latitude-time section at 83 km between 16th July and 26th August. Between about 40° N and 55° N a period-2 occurs at this height, but the amplitude decreases considerably in August. Figure 11 displays the diurnal variations of ozone at 83 km and 2.5° N under real conditions again. Next there is quasi a period-2 which jumps into a period-4 and returns finally to a period-2 again. The amplitudes of the diurnal variations changes permanently.

The water vapor concentration is the most important chemical parameter. It was found that a atmosphere that is too dry (less than 1 ppmv) or too humid (larger than about 6 ppmv) does not create a period-2. Also the ratio of daytime to nighttime hours plays an important role. In the case studied by Sonnemann and Fichtelmann (1997), the period-2 amplitude difference amplified with increasing daytime hours after a subtle start for about 2 daytime hours, but it decreased again after 14 daytime hours. The height range of nonlinear response is shifted upward if the water vapor concentration increases. That may be one reason why in high latitudes this range is located higher because the humidity also increases toward the summer pole. Another fact that has to be considered is the increase of the level of constant air density in summer and the prolongation of the sunshine hours. In

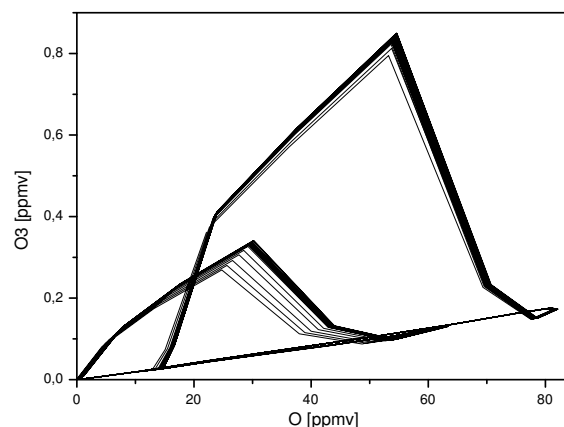


Fig. 9. Quasi O₃-O attractor at 83 km height and 42.5° N for a permanent 1st July.

the mesopause region the water vapor concentration acts in the chemical system, above all, via the product $J_{H_2O}[H_2O]$ where J_{H_2O} stands for the dissociation rate of water vapor. During the years of high solar activity J_{H_2O} is about 50% larger than the minimum value at the mesopause. Although the water vapor concentration is reduced during this phase due to enhanced photolysis, this effect does not compensate the increased dissociation rate so that the product is larger than under mean conditions anyway (Sonnemann and Grygalashvyly, 2004). A clear effect is the increase of the ozone concentration for decreasing solar activity due to the decrease of hydrogen radicals formed destroying the odd oxygen constituents ozone and atomic oxygen. Figure 12 shows a comparison of the diurnal variations of ozone at 83 km altitude and 2.5° N latitude between high (upper panel) and low solar activity (lower panel) after 30 days for a permanent 1st July. In case of high solar activity a period-2 is created, whereas for low solar activity a period-6 occurs - period doubling of a period-3. Fichtelmann and Sonnemann (1992) found that the period-3 is an inherent period of the chemical system, also indicating that the system has the potential to respond chaotically. In other regions of nonlinear response the period-2 oscillation is conserved but the amplitudes of the acting species changes according to the solar activity.

5 Discussion

5.1 Discussion of the model results

In zero-dimensional models, the height range of the nonlinear response amounts to only a few kilometers (2 or maximum 3 km). Under the action of diffusion this range theoretically extends to infinity but practically, when searching for a visible effect, only to about twice as long. An essential drawback of all three-dimensional models consists in their coarse height resolution. We use a relatively fine resolution of 1.1 km. Thus, the chemistry responds in a detectable nonlinear manner only at four or maximum at six vertical

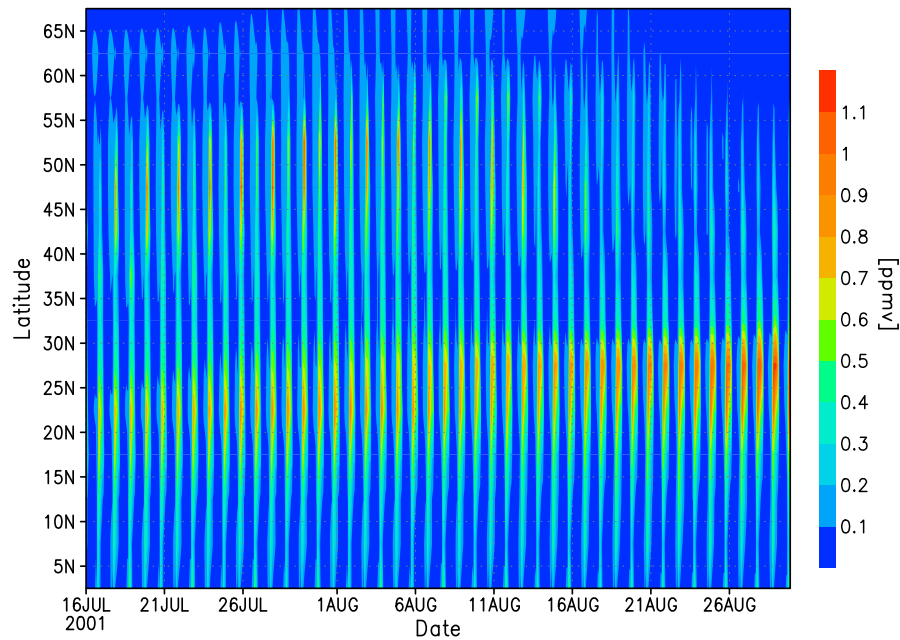


Fig. 10. Latitude-time section at 83 km between 16th July and 26th August calculated under real conditions. Between about 40° N and 55° N a period-2 occurs in this height but the amplitude decreases considerably in August.

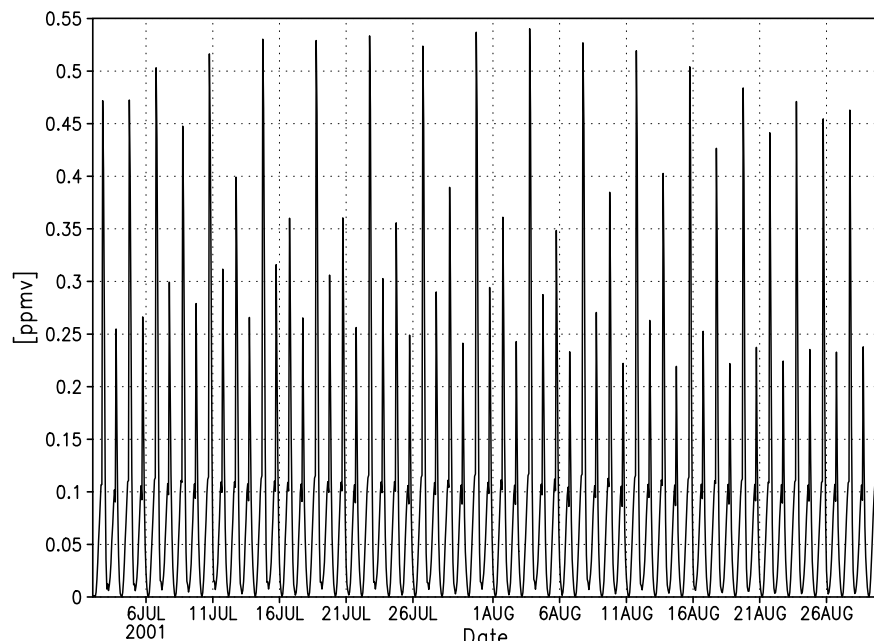


Fig. 11. Diurnal variations of ozone at 83 km and 2.5° N calculated under real conditions. There occurs a quasi period-2 oscillation in the first stage which jumps into a period-4 and returns finally to a period-2 again.

grid points. The nonlinear effects should be suppressed in all models which have a height resolution of about half a scale height. The characteristic transport time due to vertical wind is defined as that time needed to traverse the distance of a scale height. The zonally averaged vertical winds in the domain of nonlinear response have an order of few cm s^{-1} . An air parcel that moves upward or downward with a moderate constant speed of 1.736 cm s^{-1} needs two days to

cross a distance of 3 km. The essentially stronger tidal winds result in an up-and-down motion of the air parcels by approximately 1 km or a little more in high latitudes (Berger and von Zahn, 2002) but essentially more in low latitudes. The height gradient of the water vapor mixing ratio within the entire mesopause region is strong. This fact also entails, according to the vertical tidal motion, a considerable periodic diurnal and semi-diurnal variation of the water vapor

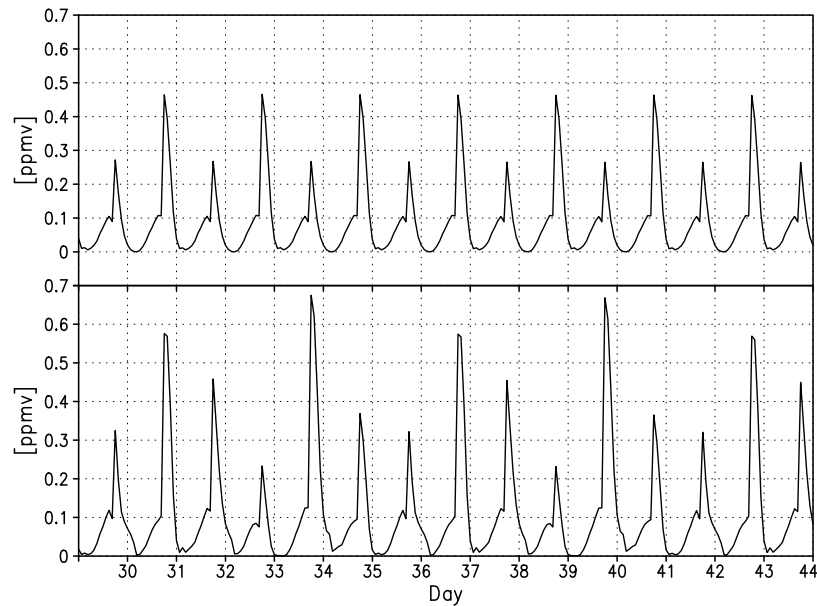


Fig. 12. Comparison of the diurnal variations of ozone at 83 km altitude and 2.5° N latitude between high (upper panel) and low solar activity (lower panel) after 30 days for a permanent 1st July. In case of high solar activity a period-2 is created whereas for low solar activity a period-6 occurs.

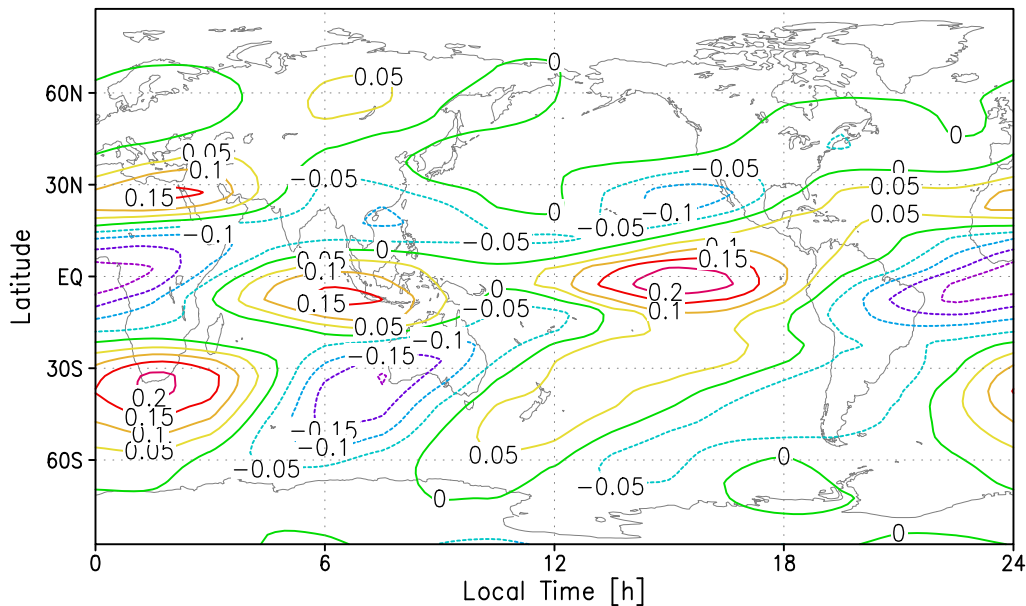


Fig. 13. Snapshot of the diurnal variation of the vertical wind in m s^{-1} around the parallels at 83 km altitude for the 1st July. (For a fixed universal time, the local time depends on the longitude.) Clearly recognizable are the patterns of the strong semi-diurnal tidal wind.

concentration in constant height, likewise acting as an enforcing mechanism. The vertical wind transports air not subjected to a nonlinear response through the domain of nonlinear response where the nonlinear response starts to influence the chemistry within the air parcel. Figure 13 exhibits a snapshot of the diurnal variation of the vertical wind in m s^{-1} around the parallels at 83 km altitude for 1st July. Clearly recognizable are the strong semi-diurnal tidal wind patterns. Obviously, if considering the results of calcula-

tions, a moderate vertical wind does not prevent the creation of a period-2 oscillation, although additionally the eddy diffusion acts to destroy the nonlinear effect. But for stronger vertical winds (clearly larger than the speed to traverse the height interval of nonlinear response, say $>2 \text{ cm s}^{-1}$) the residence time of an air parcel within the domain of nonlinear response is too short to excite a subharmonic oscillation. Consequently, stronger vertical winds can destroy the nonlinear response. We found the period-2 oscillation in regions

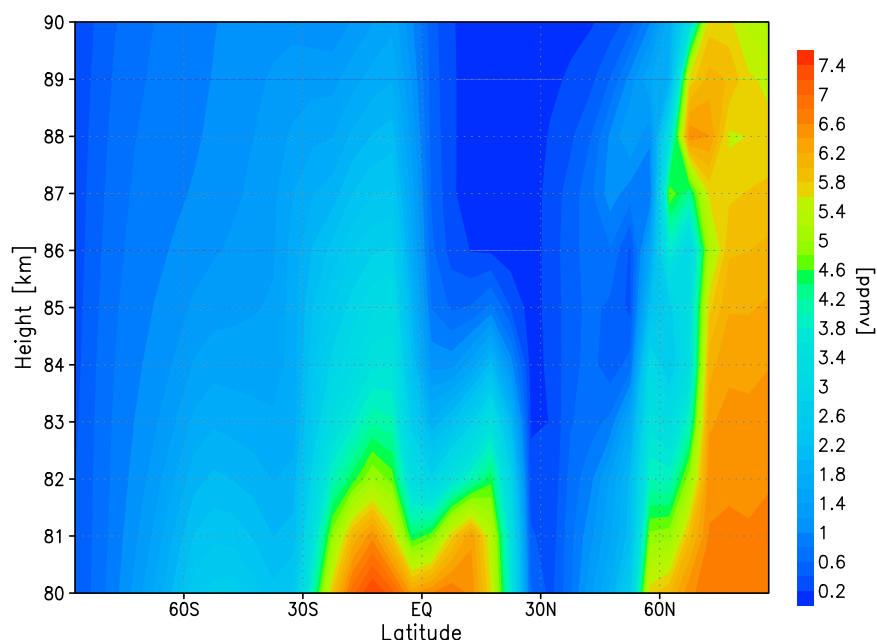


Fig. 14. Altitude-latitude section of the water vapor mixing ratio between 80 and 90 km for 1st July. The water vapor mixing ratio depends strongly on the direction and mean magnitude of the vertical wind.

of more moderate vertical winds. The calculations revealed neighboring regions marked by different behavior, but just as decisive is the transport of humid or dry air by the vertical winds from below or above. Figure 14 displays an altitude-latitude section of the water vapor mixing ratio between 80 and 90 km for 1st July again. The water vapor mixing ratio is small in those latitudinal regions of, on average, downward directed vertical winds as found around 30°N , and vice versa, it is high around the equator according to strong upward winds there. Particularly in the polar day region largest water vapor mixing ratios have been observed. They result from a permanent upward flux of humid air within the summery polar high pressure regime and from the autocatalytic water vapor formation under the action of strong solar insolation (Sonnemann et al., 2005).

The creation of subharmonic oscillations in the chemistry within the mesopause region could be a real effect in the atmosphere under favorable conditions of relatively low turbulent diffusion (the 1-D-calculations by Sonnemann and Feigin (1999a) came to a value of $<2 \times 10^5 \text{ cm}^2 \text{ s}^{-1}$), not too strong vertical wind ($< \text{about } 2 \text{ cm s}^{-1}$) and not too small water vapor concentrations (in calculations by means of simplified models the mixing ratios had to be $>1\text{--}2 \text{ ppmv}$) as those occur in the summer hemisphere. But within the region of the polar mesospheric summer echoes and noctilucent clouds (PMSE-NLC region) in high latitudes in summer a freeze drying takes place just at that height where the two-day oscillation should occur. Under this condition the water vapor mixing ratio drops to few tenth of a ppmv (Berger and von Zahn, 2002, Rapp, 2003). The model calculations, using simplified models, showed that these values are too small to create a subharmonics (e.g. Sonnemann and Fichtelmann,

1997; Feigin et al., 1998). Therefore, it is rather unlikely that the effect appears under the condition of freeze drying. The model used did not consider a micro-physical code of ice particle genesis causing the freeze dry effect. Therefore, we also got clear period-2 oscillations within the PMSE-NLC region.

5.2 Discussion of observations

Although different hints to a subharmonic period in the mesosphere and mesopause region have been published, no direct evidence could be given for a chemically induced two-day oscillation thus far. Azeem et al. (2001) investigated mesospheric ozone and temperature data obtained by the Microwave Limb Sounder (MLS) on board the UARS and found a 2-day wave in ozone most pronounced near 30 deg latitude in the southern summer hemisphere in late January. The results seem to indicate that the two-day wave in the ozone is chemically driven in the mesosphere region via changes in reaction rates that are strongly temperature dependent. However, the authors could not investigate the mesopause region, but only a region where a chemical two-day wave cannot be created.

The well-known phenomenon of the quasi two-day wave in the prevailing wind detected by Doyle in 1968 and later by Müller in 1972 is commonly interpreted in terms of dynamical processes. There is a large body of publication on this phenomenon. The same dynamical explanation is given for the two-day variation in the temperature as reported by Wu et al. (1993). The observation of a two-day variation in the oxygen green line volume emission rate by Ward et al. (1997) at a height where the characteristic time for O is too long to produce such an oscillation was interpreted by an excitation

from below. Surprisingly, the same feature with respect to the seasonal and latitudinal occurrence as found for the two-day wave in the prevailing wind have also been found for the period-2 subharmonics. The two-day wave of the prevailing wind is a typical summer effect, absent in winter and also nearly absent during the equinoxes. It is stronger in late summer than in early summer and strengthens with increasing solar activity. As is typical, the observed two-day wave in the prevailing wind within the mesopause is characterized by periods of transient quasi period-2 sometimes pulse-like oscillations and not of a permanent oscillation over the entire season. The oscillations usually remain reasonably stable over a period of several days (e.g. Poole, 1990) up to a month.

The chemical two-day oscillation induces a two-day oscillation of the chemical heating rate. As Fig. 7 indicates, the pumping by chemical heating is a robust effect. The temperature amplitude of the chemical heating rate ranges in the order necessary to excite a certain two-day wave in the prevailing wind in the mesopause region but it is too small to explain the full effect which needs minimum an amplitude of 5 K day^{-1} (Forbes, 1982). In observations, maximum amplitudes of the prevailing wind components are usually attained at heights between 80 and 95 km (e.g. Craig et al., 1980; Ward et al., 1996) which correspond to the occurrence height of the nonlinear response. However, the two-day variations of the temperature or ozone have been already found at the lower mesosphere and not only in the mesopause region (e.g. Wu et al., 1995; Azeem et al., 2001), but this could be caused by a different effect than that what creates the two-day wave in the mesopause region. The measurements of the prevailing wind indicates a maximum around 90 km altitude and only small values at 75 to 80 km (e.g. Harris, 1994; Palo and Avery, 1996; Thayaparan et al., 1997). Unfortunately, our model is not self-consistent thus far. Therefore, we cannot say whether the chemical heating essentially feeds back to the dynamics, also influencing the vertical transport above and below the region of nonlinear response. Such a transport could amplify the amplitude of the chemical heating rate due to a modulated transport of atomic oxygen (as found by Ward et al., 1997) carrying almost only the latent chemical heat. Actually relating to observations, it would be difficult to distinguish between dynamically induced or chemically caused quasi two-day variation, e.g. of measured ozone, as long as the amplitude difference between alternating days is not very different, meaning not larger than would result from a temperature induced variation. As mentioned above, the model only calculates climatological averages, meaning it does not compute variations under real short-periodic changeable conditions as they are connected with stratospheric warmings or gravity and planetary wave activity. That is why a period-2 oscillation lasts over several weeks in the model and that no pulse-like periods occur.

The period of all the observed quasi two-day waves in the prevailing wind is not exactly 48 h but varies around that value by some hours, mostly somewhat larger than 48 h. Surprisingly, the order of the difference to 48 h agrees with the

shift of the period for two days resulting from the photochemical Doppler-effect due to the zonal wind (Sonnemann, 2001). The summer east wind system results in Doppler periods larger than 48 h for two days.

There are different investigations concerning the periodicity of PMSE or NLC. A quasi 5-day variation of bright NLC has been found by Gadsden (1985). Merkel et al. (2003) confirmed this finding. In measurements of the polar mesospheric clouds (PMC – being the NLC occurring in the polar region) by means of the Student Nitric Oxide Explorer (SNOE) satellite since 1998, the 5-day wave was observed in both the northern and southern hemisphere polar summers. This variation was interpreted by Sugiyama (1994) and (1998) in terms of a limit cycle of the NLC genesis but Kirkwood et al. (2002) interpreted the 5-day wave as a planetary scale wave propagating from below. However, a two-day variation in the NLC occurrence rate has not been found. The occurrence of NLC depends sensitively on the temperature. Rosenlof and Thomas, (1990) reported about a five-day variation of mesospheric ozone between 50 and 90 km. A five-day oscillation in chemistry was found under very idealized conditions (Sonnemann and Feigin, 1999a) but it seems to be rather unlikely in realistic systems. It may be a vague hint at a higher subharmonics that Fig. 6 showed a transient period-6 but it is questionable whether or not such periods can also occur in the real atmosphere.

The chemistry in the highest summery latitudes is not subjected to any nonlinear effects in our model. We found still a two-day oscillation at 77.5° N but not more at 82.5° N . The main reason certainly consists in the fact that there is no or only a very short night, meaning only a smoothed diurnal variation which could drive the chemical oscillator (Sonnemann and Fichtelmann, 1997). The photochemical system has an unequivocal solution for all sets of constant dissociation rates. As the system is a nonlinear radiatively driven chemical oscillator, it needs a radiative excitation by a clear day-to-night difference. In winter additionally, the water vapor concentrations are too small to create a period-2 oscillation.

6 Summary and conclusions

The main goal of this paper was to demonstrate that the chemical system in the mesopause region is able to create subharmonic oscillations under nearly realistic conditions. The global 3-D-model of the dynamics and chemistry produces these phenomena within the mesopause region and within certain latitudinal regions, particularly during the summer season under the condition of relatively weak vertical winds, not too strong turbulent diffusion, and not too small water vapor concentrations. The occurrence of transient subharmonics after a sudden change of a system parameter, such as turbulence, vertical wind velocity or humidity, seems to be possible. The two-day oscillation of the active chemical constituents entails a two-day variation of the chemical heating rate which feeds back to the dynamics. As

our model does not operate interactively, this response cannot be calculated by means of our present model version. Therefore, the development of a self-consistent model in which the chemical heating rate also feeds back to the dynamics is the first priority. Different measurements in the upper mesosphere/mesopause region revealed a periodicity of quasi two days, however, there is no direct observational evidence thus far for a real chemical period-2 oscillation caused by the non-linearity of the photochemical system.

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